# On the origin of exponential growth in induced earthquakes in Groningen

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**Abstract.** The Groningen gas field shows exponential growth in earthquake event counts around a magnitude M1 with a doubling time of 6-9 years since 2001. This behavior is identified with dimensionless curvature in land subsidence, which has been evolving at a constant rate over the last few decades essentially uncorrelated to gas production. We demonstrate our mechanism by a tabletop crack formation experiment. The observed skewed distribution of event magnitudes is matched by that of maxima of event clusters with a normal distribution. It predicts about one event < M5 per day in 2025, pointing to increasing stress to human living conditions.

Keywords: Induced earthquakes; crack formation; statistical forecasting

## 1. Introduction

Induced earthquakes are a modern phenomenon associated with exploitation of natural energy resources or injection of waste (Rayleigh *et al.* 1976, Nicholson and Wesson 1990, Frolich 2012, Evans *et al.* 2012) at generally moderate strengths (e.g., Ellsworth 2013). Recent studies in particular focus on the potential for a correlation between induced events and fracking and waste water injection (Keranen *et al.* 2014, Weingarten *et al.* 2015). Fig. 1 shows induced events in Groningen, The Netherlands, due to the exploitation of natural gas since 1959 (De Kam 2014).

Though moderate in magnitude, frequent events act as physical stressors to living conditions. As a relatively modern phenomenon, induced events pose some novel challenges for hazard analysis, beyond a mere extrapolation of the geophysical impact of natural earthquakes (Hough 2013). In

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Groningen, half of the inhabitants are now living with a feeling of anxiety or fear and wish to move out, in part, by the stochastic nature of earthquakes (Fonteyne *et al.* 2009, Grillon *et al.* 2004). It has led to ameliorating policies (e.g., De Waard (2014)), generally seeking reduction in earthquake activity by limiting gas production, whose short term impact is uncertain (Nicholson and Wesson 1990, Bakun *et al.* 2005, NAM 2013).

Groningen data on reservoir pressure, land subsidence and yearly production of natural gas cover an area of 900 km<sup>2</sup> (NAM 2013). Since 2001, earthquake event counts show exponential growth at seemingly 'clock-work' precision (Fig. 1), based on a logarithmic plot of event counts with remarkably good fit by a linear curve covering the year 2000 to the present. This observation applies to event counts with a choice of magnitude cut-off over a broad range  $M_c = 0-3$ . Collectively, they represent an energy release  $\dot{E} \simeq 3$ GJ in over one hundred registered events per year with a doubling time of 6-9 years (see also Dost *et al.* (2013)). The Groningen field apparently shows a robust principle at work with unknown but possibly long-term implications by exponential growth in counts.

Fig. 2 shows NAM production levels in GNm<sup>3</sup> per year and land subsidence, associated with a decline in reservoir pressure to below 100 bar at present, down from about 150 bar in 2000 and 350 bar in 1959. The statistical correlation by Pearson coefficient between event counts in Fig. 1 to production is nill, indicative of no correlation (Fig. 2, lower left panel).

These earthquakes arise from stresses induced by a decline in well pressure (Hubbert and Rubey 1959), possibly at some distance away and delayed in time. Paradox Valley in Colorado (Ellsworth 2013) shows events at distances over ten kilometers with delays of a decade. The Salton Sea Geothermal Field shows, subsequent to initial exploitation, a correlation of seismicity with net extracted volume  $\dot{V}$  (injected minus extracted volume) of about  $10^6$  m<sup>3</sup> per month at 1-2.5 km depth (Brodsky and Laipote 2013). Large area induced seismicity points to large scale deformations that may be measured in land subsidence.

Here, we propose a novel physical mechanism for exponential growth in event counts in induced earthquakes in response to dimensionless curvature of land subsidence, demonstrated by a table top experiment. Based on dimensionless variables, it is insensitive to details of geophysical composition. The skewed distribution in magnitudes is modeled by maxima of event clusters, that will be used for a statistical forecast on event count and maximal magnitude in the coming decade.

# 2. Exponential growth from curvature induced strain

Groningen land subsidence is  $\dot{h} \simeq 0.5$  cm yr<sup>-1</sup> (Fig. 2, right lower panel) with a delay in response to a reduction in volume V due to a gradual compactification in the reservoir associated with a steady decrease in reservoir pressure (NAM 2013). Compactification  $\dot{V} \simeq 10^6$  m<sup>-3</sup> per month is roughly similar to that in aforementioned Salton Sea Geothermal Field. It represents a mechanical energy input of  $\dot{E}_0 \simeq 0.3$ TJ per year, whereby  $\dot{E}/\dot{E}_0 \simeq 10^{-2}$ .

Subsidence gives rise to curvature-induced strain, inevitably leading to the build up of internal stresses. At critical values, materials fail by cracks or slip, described by the Mohr-Coulomb law (Coulomb 1776, Hubbert and Rubey 1959). Subsidence in Groningen is taken the shape of a shallow bowl of radius  $r \simeq 15$  km (NAM (2013), Fig. 3). Relative to its radius of curvature R, the dimensionless curvature  $\rho = r/R \simeq 4 \times 10^{-5}$  should be representative for large scale internal strain. For small perturbations in the absence of cracks and slides, internal stresses increase with strain induced

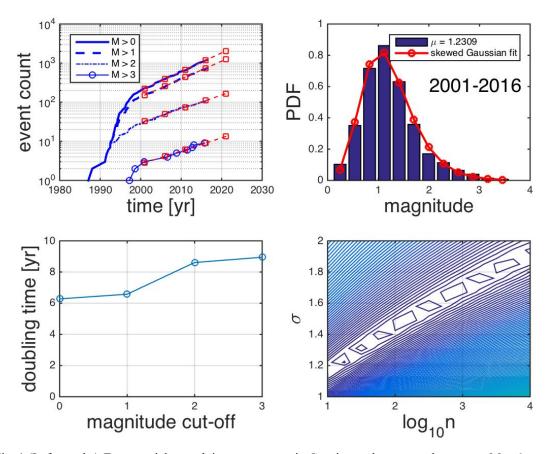


Fig. 1 (Left panels.) Exponential growth in event counts in Groningen that extrapolate to one M > 0 event per day in 2025. (Right panels.) The distribution of magnitudes is skewed, here modeled by that of maxima of event clusters of size n, drawn from a normal distribution with dispersion  $\sigma$ . A contour plot of residual least square ( $L_2$ ) errors (bottom right) shows best-fit parameters  $n \simeq 2250$ ,  $\sigma \simeq 1.7$  (white strip)

at a well, satisfying a linear elliptical equation subject to free surface conditions on the land above and maximal strain and pressure gradients in the well at the reservoir depth of about 3 km. Following Fig. 3, internal strain hereby scales linearly with subsidence  $\delta h = R(1 - \cos \theta) \simeq R\theta^2/2$ ,  $R \sin \theta = r$ , whereby  $\delta h \simeq r^2/2R$ , i.e.

$$\frac{\delta h}{r} \simeq \frac{1}{2}\rho. \tag{1}$$

Glitches appear by cracks and slides when internal strains  $\epsilon$  reach a critical value  $\epsilon_c$  with associated critical values for (1). Such impulsive release of internal stress is followed by relaxation to sub-critical values of  $\epsilon$ .

Glitches partition a region into two followed by a relaxation  $\eta$  to a sub-critical state, satisfying

$$0 < \eta < \epsilon_c. \tag{2}$$

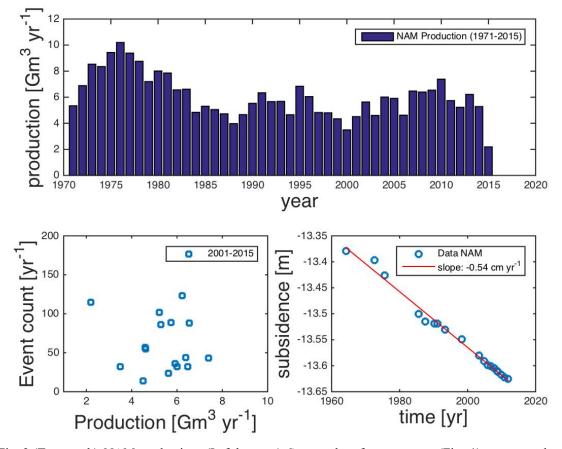


Fig. 2 (Top panel.) NAM production. (Left bottom.) Scatter plot of event counts (Fig. 1) versus production over 2001-2015. The absence of a correlation (Pearson coefficient 0.0816) evidencies a lack of sensitivity to NAM production. (Right bottom.) Subsidence increases essentially linear in time since 1970. (Data from NAM (2013))

Continuing subsidence  $\delta h/r < 0$  increases  $\eta$ , back to  $\epsilon_c$  over time. In approaching  $\epsilon_c$ , regions in the new partition crack or slice into two once again. A time-evolving strain  $\epsilon = \epsilon(t)$  is hereby restricted to the range (Fig. 3)

$$\epsilon_c - \eta < \epsilon < \epsilon_c. \tag{3}$$

As  $\rho$  continues to grow by subsidence, a shallow bowl forms in the region around the natural gas reservoir. Induced strain  $\epsilon$  hereby extends approximately uniformly over the entire area of the bowl and (3) produce exponential growth in event count

$$N = N_0 e^{\alpha \rho},\tag{4}$$

where  $\alpha$  is a dimensionless constant. For Groningen,  $N \simeq 10^3$  today (since 2001),  $\dot{N}/N \simeq 0.1$ ,  $\rho \simeq 4 \times 10^{-5}$  and  $\dot{\rho} \simeq 6.7 \times 10^{-6}$ , give the estimates  $\alpha \rho \simeq 6$  and  $N_0$  is of order a few, as expected.

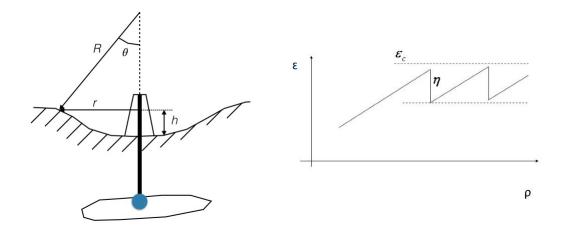


Fig. 3 (Left.) Geometry of land subsidence around a well. Internal strain  $\epsilon$  scales linearly with dimensionless curvature  $\rho = r/R$  due to land subsidence. (Right.) Evolution of  $\epsilon$  as a function of  $\rho$  to critical values  $\epsilon_c$  with glitches and relaxation by  $\eta$  to sub-critical values

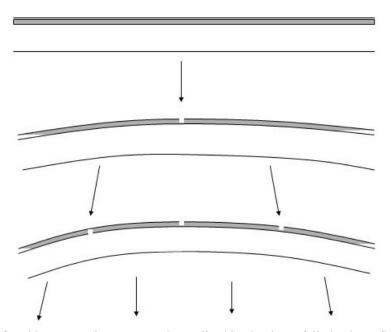


Fig. 4 Schematic of a tabletop experiment on cracks, realized in aluminum foil glued to a flexible substrate of length  $L \simeq 20$  cm subject to bending. The first crack induced by dimensionless curvature L/R defined by the radius of curvature R forms at a critical strain  $\epsilon = \epsilon_c$ . Each such partition produces two sections, followed by relaxation to a stressed but stable state at a strain  $\epsilon_c - \eta$ . Further bending rebuilds stress back to  $\epsilon_c$ , causing both sections to crack, again followed by relaxation

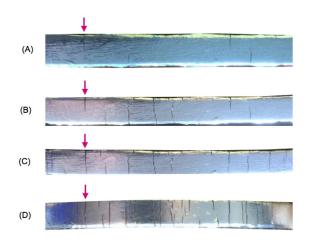


Fig. 5 Snapshots of crack formation in the aluminum foil subject to bending, increasing from A-D. Arrows point to a single crack for reference

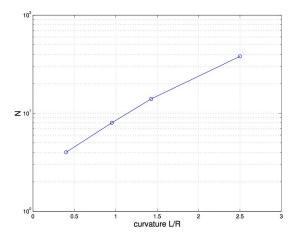


Fig. 6 The observed number of cracks N shows an approximately exponential growth with  $\rho$ 

Figs. 4-6 shows a table top experiment demonstrating (4). This  $\epsilon$  excess of  $\epsilon_c$  is visible in cracks in aluminum foil, here glued to a flexible substrate as a function of dimensionless curvature  $\rho$ . In this experiment,  $\eta$  is essentially equal to the critical value  $\epsilon_c$ . It serves to illustrate that, in general,  $\eta$  need not be small and can be on the order of  $\epsilon_c$ .

Our mechanism (4) is completely dimensionless. It does not explicitly depend on the details of internal length, area or volume or geophysical structure. Independent of dimension, it has implications also to three dimensional microfracturing experiments (Scholz 1968) on induced fractures. This applies especially, since the latter are typically limited to a few hundred meters (Davies *et al.* 2012, Lacazette and Geiser 2013), which is rather local relative to the scale depth of about 3 km of the Groningen reservoir. Given that subsidence in the Groningen field increases at an approximately constant rate in time (Fig. 2), we therefore identify the exponential trend (Fig. 1) with (4).

Our model does not address time delay between exploitation and evolution of subsidence. Time

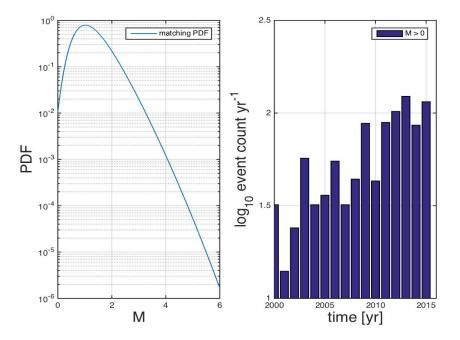


Fig. 7 Overview of a statistical forecast based on the optimal fit of the PDF in (5) to the observed skewed distribution of events in Fig. 1

delays are notoriously varied in different locations, and their origin remains somewhat mysterious especially so when delays are long. For Groningen, the lack of any correlation between gas exploitation and observed event counts (Fig. 2) over the course of one decade is remarkable. It points to long-lived delayed events, likely on the order of a decade.

## 3. Parameter-free statistical forecast of magnitude limit

For Groningen, we attribute induced events to sudden release of energy in clusters of essentially contemporaneous events, derived from normal distributions in  $\eta$  and  $\epsilon_c$ . The magnitude of an earthquake (as reported) is then the maximum of the  $\eta$  in such cluster (cf. the composite source model of (Frankel 1991)).

In general, a collection of n independent samples of a normal variable x with mean zero and dispersion  $\sigma$  produces a truncated Gaussian distribution, whenever n is finite, and maxima thereof satisfy a skewed distribution. To see this, the probability density function (PDF) of the absolute values satisfies (e.g., van Putten *et al.* 2014, van Putten 2016)

$$p(x) = \frac{n}{\sigma} \sqrt{\frac{2}{\pi}} \operatorname{erf}(y)^{n-1} e^{-y^2}$$
(5)

with  $y = x/\sqrt{2}\sigma$ , derived from the probability  $P(\langle x) = \int_0^x p(s)ds$  that all *n* samples are less than |x|. Fig. 1 shows that (5) accurately matches in the observed distribution of Groningen for n = 2250

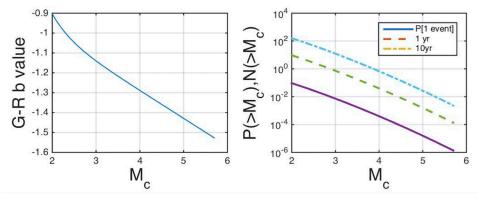


Fig. 8 The PDF in (5) to the observed skewed distribution of events in Fig. 1 has an associated Gutenberg-Richter *b*-value -1.1 < b < -1.3 ( $3 < M_c < 4$ ) (left). Extrapolation of the current yearly event rates (right top) and the cumulative probability  $P(> M_c)$  derived from the fit (5) shows  $N(> M4) \simeq 1$  and  $P(> M5) \simeq 1\%$  over the next decade (right)

with  $\sigma = 1.7$ , which is an almost degenerate minimum along a line in the  $(n, \sigma)$  plane. Following Fig. 3, the observed magnitudes are the maxima in  $\eta$  associated with the random variable  $x = \epsilon_c$ . Our match to the distribution of the Groningen events involves a shift downwards from the mean  $\mu = 6.4$  in x.

Figs. 7-8 shows that our matching PDF derived from (4) has a decay that is slightly steeper than exponential with associated slowly varying Gutenberg-Richter *b* value 1.1-1.3, consistent with earlier estimates that depend on a choice of magnitude cut-off  $M_c$  (Ishimoto and Lida 1939, Gutenberg and Richter 1944, 1954, Schorlemmer *et al.* 2005). For statistical forecasting, our PDF takes into account weak deviations from exponential decay and obviates the need for using a choice of magnitude cut-off.

#### 4. Summary and outlook

Exponential growth in Groningen event counts is a remarkable man-made phenomenon. We attribute it to curvature-induced strain in light of the relatively constant rate of land subsidence observed over the last few decades. Based on dimensionless variables and demonstrated by experiment, our mechanism is robust and does not depend on the detailed geophysical struture of the Groningen soil composition.

Fig. 1 shows a doubling time of 6-9 years in total event count and the observed PDF of event magnitudes with positive skewness. Figs. 7-8 show a matching PDF of maxima of event clusters (5). This matching PDF shows a Gutenberg-Richter value that, in absolute value, increases steeply with the cut-off  $M_c$  (Fig. 8). While it assumes typical values  $|b| \simeq 1.15 - 1.3$  for  $M_c \simeq 3 - 4$ , |b| increases to about 1.5 for  $M_c \simeq 5 - 6$ . As a result, we arrive at the following forecast for 2025:

- An event rate N(>M0) of one per day;
- $N(>M4) \simeq 1$  per year;
- A probability of 1% for one >M5 event per year.

The relatively small probability of the latter is due to aforementioned increase in |b|-value in the tail of the PDF, larger than typical estimates based on moderate cut-off magnitudes  $M_c$ .

Based on  $\dot{E}/\dot{E}_0 \simeq 10^{-2}$ , there is no energetic limit to the observed exponential growth for decades to come. (Theoretically, it can continue for another half a century.) An apparent slow large scale creep delayed relative to production is consistent with a lack of a correlation of yearly subsidence with fluctuations in gas production and reservoir pressure, whose physical origin, however, remains somewhat mysterious. It suggests a significant time delay time on the order of a decade or more, that challenge the formulation of effective policies and strategies to ameliorate this bleak outlook as a well-populated area (cf. (van Eck *et al.* 2006, Peterson *et al.* 2008)).

While the number of earthquakes is rapidly increasing with the potential for significant continuing stresses to human living conditions, their strength probably remains below M5. For Groningen, we propose that the exponential growth in event counts is included in hazard analysis as a general stressor to living conditions. Also, the relatively high frequency spectrum of induced motions is potentially relevant, as it falls in the range of eigenfrequencies of around 0.1-25 Hz, where the higher frequencies are expected to be relevant to homes in the Groningen area. (High frequency earthquakes may be found, for instance, in the spectrum of induced events at the Hellisheidi Geothermal Powerplang in Iceland (Halldorsson *et al.* 2012) peaks at 10 Hz.) These relatively high frequencies *f* are commonly attributed to the scaling  $f \propto M^{-1/3}$  with magnitude (Akii 1967, Bune 1970, Kanamori and Anderson 1975, Anderson 2003, Aki and Richards 1980).

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#### References

Aki, K. (1967), "Scaling law of seismic spectra", J. Geophys. Res., 72(4), 1217-1231.

- Aki, K. and Richards, P.G.(1980), *Quantitative Seismology, Theory and Methods*, W.H. Freeman and Co. San Francisco, CA, USA.
- Anderson, J.G. (2003), in Int. Handbook of Earthquake & Engineering Seismology, Part 2, W.H.K. Lee, H. Kanamori, P. Jennings an C. Kisslinger (eds.), Ch. 57, p.937.
- Bakun, W.H., Aagaard, B., Dost, B., Ellsworth, W.L., Hardebeck, J.L., Harris, R.A. and Michael, A.J. (2005), "Implications for prediction and hazard assessment from the 2004 Parkfield earthquake", *Nature*, 437(7061), 969-974.
- Brodsky, E.E. and Lajpote, L.J. (2013), "Salton sea geothermal field anthropogenic seismicity rates and operational parameters at the salton sea geothermal field", *Science*, 341(6145), 543-546.
- Bune, J.N. (1970), "Tectonic stress and the spectra of seismic shear waves from earthquakes", J. Geophys. Res., 75(26), 4997-5009.

- Coulomb, C.A. (1776), "Essai sur une application des règles des maximis et minimis à quelques problàmes de statique relatifs à l'architecture", *Mem. Acad. Roy. Div. Sav.*, 7, 343-387.
- Davies, R.J., Mathias, S.A., Noss, J., Hustof, S. and Newport, L. (2012), "Hydraulic fractures: How far can they go?", *Marine and Petrol. Geol.*, 37(1), 1-6.
- De Kam, G. (2014), "Fear of earthquakes risen sharply among inhabitants of earthquake zone", http://www.rug.nl/news-and-events/news/archief2014/nieuwsberichten/angst-voor-aardbevingen-onderinwoners-aardbevingsgebied-sterk-gestegen?lang=en.
- De Waard, P. (2014), De Volkskrant, January 14.
- Dost, B., Caccavale, M., van Eck, T. and Kraaijpoel, D. (2013), "KNMI Report on expected PGV and PGA values for induced earthquakes", KNMI, The Netherlands.
- Grillon, C., Baas, J.P., Lissek, S., Smith, K. and Milstein, J. (2004), "Anxious responses to predictable and unpredictable aversive events", *Behav. Neuroscience*, **118**(5), 916-924.
- Gutenberg, B. and Richter, C.F. (1944), "Frequency of earthquakes in California", *Bull. Seismol. Soc. Am.*, **34**(4), 185-188.
- Gutenberg, B. and Richter, C.F. (1954), Seismicity of the Earth and Associated Phenomena, Princeton University Press, Princeton, NJ, USA.
- Ellsworth, W.L. (2013), "Injection-induced earthquakes", Science, 341(6142), 142-147.
- Evans, K.F., Zappone, A., Kraft, T., Deichmann, N. and Moia, F. (2012), "A survey of the induced seismic responses to fluid injection in geothermal and CO2 reservoirs in Europe", *Geothermics*, **41**, 30-54.
- Fonteyne, R., Vervliet, B., Hermans, D., Baeyens, F. and Vansteenwegen, D. (2009), "Reducing chronic anxiety by making the threatening event predictable: An experimental approach", *Behav. Res. and Therapy*, **47**(10), 830-839.
- Frankel, A.J. (1991), "High-frequency spectral falloff of earthquakes, fractal dimension of complex rupture, b value, and the scaling of strength on faults", *Geophys. Res.*, **96**, 6291-6302.
- Frohlich, C. (2012), "Two-year survey comparing earthquake activity and injection-well locations in the Barnett Shale, Texas", *PNAS*, **109**(35), 12934-12938.
- Halldorsson, B., Ólafsson, S., Snaebjörnsson, J.Th., Sigurosson, S.U., Rupakhety, R. and Sigbjörnsson, R. (2012), "On the effects of induced earthquakes due to fluid injection at hellisheidi geothermal power plant, Iceland", *Proceeding of the 15th WCEE*, Lissabon, Portugal.
- Hough, S.E. (2013), "Spatial variability of "Did you feel it" intensity data: insights into sampling biases in historical earthquake intensity distributions", *Bull. Seism. Soc. Am.*, **103**(5), 2767-2781.
- Hubbert, M.K. and Rubey, W.W. (1959), "Role of fluid pressure in mechanics of overthrust faulting I. Mechanics of fluid-filled porous solids and its application to overthrust faulting", *Geol. Surv. Bull.*, **70**(2), 115-166.
- Ishimoto, M. and Lida, K. (1939), "Observations sur les seismes enregistres par le microsismographe construit dermement", Bull. Earthquake Res. Inst., Tokyo Univ., 17, 443-478.
- Kanamori, H. and Anderson, L. (1975), "Theoretical basis of some empirical relations in seismology", Bull. Seismol. Soc. Am., 65(5), 1073-1095.
- Keranen, K.M., Weingarten, M., Abers, G.A., Bekins, B.A. and Ge, S. (2014), "Sharp increase in central Oklahoma seismicity since 2008 induced by massive wastewater injection", *Science*, **345**(6195), 448-451.
- Lacazette, A. and Geiser, P. (2013), "A new method of neostress determination from passive seismic data", *Marine and Petrolium Geology*, 43, 516-518.
- Nederlandse Aardolie Maatschappij (2013), "Technical Addendum to the Winningsplan Groningen, Subsidence, Induced Earthquakes and Seismic Hazard Analysis in the Groningen Field," http://www.rijksoverheid.nl/documenten-en-publicaties/rapporten/2014/01/17/bijlage-1-analyse-over-

verzakkingen-geinduceerde-aardbevingen-en-seismische-risico-s.html, NAM, The Netherlands.

- Nicholson, C. and Wesson, R.L. (1990), "Earthquake hazard associated with deep well injection", US Geol. Surv. Bull., 1951 pubs.usgs.gov/bul/1951/report.pdf.
- Peterson, M.D., Frankel, A.D., Harmsen, S.C., Mueller, C.S., Haller, K.M. and Wheeler, R.L., et al. (2008), US Geol. Surv. Open-File Rep., 2008-1128.
- Rayleigh, C.B., Healy, J.H. and Bredehoeft, J.D. (1976), "An experiment in earthquake control at Rangely, Colorado", *Science*, **191**, 1230-1237.
- Scholz, C.H. (1968), "The frequency-magnitude relation of microfracturing in rock and its relation to earthquakes", Bull. Seismol. Soc. Am., 58(1), 399-415.
- Schorlemmer, D., Wiemer, S. and Wyss, M. (2005), "Variations in earthquake-size distribution across different stress regimes", *Nature*, 437(7058), 539-542.
- van Eck, T., Goutbeek, F., Haak, H. and Dost, B. (2006), "Seismic hazard due to small-magnitude, shallow-source, induced earthquakes in The Netherlands," *Eng. Geol.*, **87**(1), 105-121.
- van Putten, M.H.P.M., Guidorzi, C. and Frontera, F. (2016), "Broadband turbulent spectra in gamma-ray burst light curves", *Ap. J.*, **786**(2), 146-154.
- van Putten, M.H.P.M. (2016), "Directed searches for broadband extended gravitational-wave emission in nearby energetic core-collapse supernovae", *Ap. J.*, **819**(2), 169-181.
- Weingarten, M., Ge, S., Godt, J.W., Bekins, B.A. and Rubinstein, J.L. (2015), "High-rate injection is associated with the increase in U.S. mid-continent seismicity", *Science*, **348**(6241), 1336-1340.

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